Carbon isotopes as a sequence stratigraphic tool: examples from the Lower and Middle Toarcian marly limestones of Portugal

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ABSTRACT

The Lower and Middle Toarcian series of the Lusitanian Basin (Portugal) are generally characterized by hemipelagic marl-limestone alternations, included in the S. Gião Formation. The succession is informally organized into three 3rd order depositional sequences (ST1 to ST3), each bounded by regional discontinuities. These depositional sequences are characterized in terms of vertical and lateral litho- and biofacies changes. Using accurate ammonite biostratigraphic control, sequence boundaries are assigned to the lowermost polymorphum Zone (SBT1), polymorphum-levisoni Zone boundary (SBT2), intra-bifrons Zone (SBT3) and uppermost bonarelli Zone (SBT4). Maximum flooding intervals are identified in the uppermost polymorphum (ST1), uppermost levisoni (ST2) and uppermost bifrons (ST3) Zones. The uppermost part of levisoni Zone is a 2nd-order maximum transgressive phase. Carbon isotope data from bulk carbonate samples of selected sections in the Lusitanian Basin (Rabaçal, Porto de Mós, Coimbra and Figueira da Foz) reveal this analytical technique as good indicator of 3rd order sequence evolution. Positive δ¹³C excursions correspond to transgressive (transgressive systems tracts) phases whereas negative trends/excursions are related to periods of regression (lowstand and highstand systems tracts). Furthermore, the global δ¹³C events recognised in the Early Toarcian sections of several North-European and Tethyan basins are observed in all the studied sections of Lusitanian Basin.

Key words: Carbon isotopes, marl-limestone alternations, Portugal, stratigraphic sequences, Toarcian.

Introduction

In Portugal, Early Jurassic age sediments crop out in two different geo-structural settings, both bounded by the Iberian Variscan Massif: the Algarve Basin, in the south of Portugal and in the Lusitanian Basin, located in west-central Portugal. In the Lusitanian Basin (LB), the Toarcian is generally dominated by marl/argillace-
ous limestone and limestone alternations, usually characterised by a rich and diverse nektonic and benthic macrofauna (Duarte, 1997; Duarte et al., 2001, 2004a).

The hemipelagic nature of the Toarcian sediments in the LB and the accurate biostratigraphic control, constitute a good context for carbon and oxygen isotope studies. This kind of pelagic sediment is suitable for using isotopes as palaeoceanographic/palaeoclimatic indicators (e.g. Schölle and Arthur, 1980; Weissert, 1989; Follmi et al., 1994; Weissert et al., 1998; Kump and Arthur, 1999; Jenkyns et al., 2002). Carbon isotopes have proven a powerful stratigraphic tool in the study of the Toarcian of some European and Tethyan regions (Jenkyns and Clayton, 1986, 1997; Jiménez et al., 1996; Röhl et al., 2001; Rosales et al., 2001; Jenkyns et al., 2002; Schmid-Röhl et al., 2002; Mattioli et al., 2004, among others).

The record of stable isotopes in the Toarcian sediments of the LB also proves this parameter as a useful tool in sequence stratigraphy analysis, in particular for the definition of 3rd order sequences. Thus, the aim of this paper is to characterize the depositional systems in the whole LB during the Early and Middle Toarcian in terms of sequence stratigraphy and carbon isotope data, improving and extending some isotopic data before published in Duarte (1998).

**Geological context and lithostratigraphy**

The studied successions of Toarcian age are included in the Late Triassic (Norian?)-Callovian cycle of Wilson et al. (1989). In this cycle, the Pliensbachian and Toarcian sediments correspond to the maximum transgressive facies interval (Soares et al., 1993).

Throughout the basin, Toarcian hemipelagic deposits contain a rich nektonic (ammonites and belemnites) and benthic (mainly brachiopods, bivalves, crinoids and siliceous sponges) macrofauna. The integration of several stratigraphic/sedimentological analyses (e.g. lithofacies analysis, microfacies, sequential evolution and palaeontological evolution) indicates that deposition occurred on a carbonate ramp (homoclinal ramp sensu Read, 1982, 1985) dipping towards the northwest (Duarte 1997; Duarte et al., 2001). In the LB the Toarcian is well exposed, including the two likely depth extremes of the basin: Figueira da Foz-Cantanhede at the north end and Arrábida-Sesimbra at the south end (Figure 1).

**Stratigraphic Chart**

Excluding the Arrábida-Sesimbra, Tomar and Peniche sectors (Figure 1 and 2), the Toarcian successions of LB are very rich in ammonites and calcareous nanofossils, allowing good biostratigraphic control of the series (Elmi et al., 1989; Rocha et al., 1987, 1996; Perilli and Duarte, 2006). The studied succession belongs to the S. Gião Formation (Duarte and Soares, 2002), a marl dominated unit that ranges in age from the Early Toarcian (polymorphum Zone) up to the Late Toarcian (base of meneghinni
Zone). This unit occurs between the Leme de (Late Pliensbachian) and Póvoa da Lomba (Late Toarcian to Early Bajocian) Formations.

The S. Gião Formation can be subdivided into five informal members (Figures 2 and 3), easily recognizable across a large area of the LB (Duarte and Soares, 2002): Marly limestones with *Leptaena* facies Member (MLLF), Thin Nodular Limestones Member (TNL), Marls and marly limestones with Hildaites and Hildoceras Member (MMLHH), Marls and marly limestones with sponge bioconstructions Member (MMLSB) and Marls and marly limestones with brachiopods Member (MMLB). Generally, the thickness of each unit increases from southeast (Tomar region) towards the northwest (Figueira da Foz sector), following the direction of the dipping ramp. With the exception of Arrábida-Sesimbra sector, the ammonite records suggest that no important sedimentary gap is present in the Toarcian series of the LB.

**Methods**

**Studied Sections**

The main sections of Toarcian in the LB have been previously studied, bed-by-bed, in terms of an integrated stratigraphic analysis and sedimentary characterization (Duarte, 1997; Duarte et al., 2001, 2004a). In terms of geochemical analysis four sections were selected on the basis of relative vertical continuity, accurate biostratigraphic control, and palaeogeographic context, all cropping out in northern and central part of the basin where the Toarcian is thickest and developed in hemipelagic facies (Figures 1 and 3).
Rabaçal Section: This constitutes the reference section for the Toarcian marly limestone facies of the LB due to its excellent exposure and relatively intensive investigation (Mouterde et al., 1964/65; Duarte 1995, 1998, 2004; Duarte and Soares, 2002, and references therein). Rabaçal constitutes the type-locality of S. Gião Formation.

Porto de Mós composite-section: This is composed of the Porto de Mós I and Penas do Castelo sub-sections (Duarte, 1995, 1998). Only the middle part of the MMLHH (base of bifrons Zone; Mouterde and Ruget, 1967) is not sampled.

Coimbra composite-section: In this area, the Toarcian deposits are widespread and the succession expanded, showing well-developed hemipelagic characteristics (Duarte et al., 2004a). This section is composed by the following five sub-sections: Ribeiro, Fornos, Marmeleira, Cabeço da Azevêda and S. Simão (Duarte, 1995). Only part of MMLHH/MMLSB boundary (part of bifrons Zone) and the middle part of MMLSB (top of gradata Zone) have not been studied and sampled.

Figueira da Foz composite-section: This constitutes the least studied sector of the LB, composed by
Vale das Fontes and Brenha sub-sections (Duarte, 1995). Exposure conditions restrict observations and sampling to the MLLF, TNL and middle part of MMLHH.

Carbon and Oxygen Isotopes

To investigate $\delta^{13}C$ and $\delta^{18}O$, 109 samples of bulk-rock micritic carbonate sediments from calcareous levels (with more than 60% of carbonate) belonging to the Lower and Middle Toarcian units of Coimbra and Figueira da Foz (the two most distal sectors of LB) were analysed. Although belemnites and brachiopods are particularly rich in some units of the Toarcian, they are not present in the whole succession, which prevents construction of a continuous isotopic curve based upon these particular taxa.

All rock samples were pulverized in porcelain dish to obtain particles smaller than 80 mesh. The oxygen and carbon isotopic values were obtained from individual samples using a Kiel Carbonate Device III coupled to a MAT 252 ThermoFinnigan mass spectrometer at Petrobras Research Center (Cenpes, Brazil). This process uses approximately 10 mg of sample, which was placed in a vacuum chamber and reacted with phosphoric acid for 6 minutes at 70°C. The water and carbon dioxide generated were captured at -170°C with liquid nitrogen. Then, this mixture was brought to -110°C when the subsequently released pure CO$_2$ was again collected at -170°C with liquid nitrogen. All results are reported in parts per mil (‰) notation ($\delta$) relative to Peedee Belemnite international standard (PDB; Craig, 1957) and calibrated by routine preparation and analysis of the carbonate standard NBS-19 ($\delta^{13}C=1.95‰$, $\delta^{18}O=-2.20‰$). Day to day precision ($\pm 1\delta$) based on NBS-19 analyses was $\pm 0.05$ for $\delta^{13}C$ and $\pm 0.08$ for $\delta^{18}O$.

We also include in this work other available isotope data previously published in Duarte (1998) from Rabaçal (22 samples) and Porto de Mós (24 samples) sections.

Results: Carbon and oxygen isotope stratigraphy

Carbon Isotopes

In the newly studied sections, the $\delta^{13}C$ data, presented in the Tables I and II, range between -0.87‰ and +4.02‰, results that are consistent with normal values for marine carbonate sediments. However, significant oscillations are observed between each stratigraphic unit, such as zone, member or depo-
tional sequence, supporting the use of $\delta^{13}$C as a stratigraphic tool in the Lower and Middle Toarcian studies of the LB.

Lowest values of $\delta^{13}$C are observed at the base of levisoni Zone (TNL), where the $\delta^{13}$C signal varies between -0.87‰ in Coimbra and -0.25‰ in Figueira da Foz (1.08‰ in Rabaçal and 0.20‰ in Porto de Mós). Highest values are observed in the middle part of levisoni Zone, in the base of MMLHH, showing values above 4‰, with 4.02‰ in Coimbra (4.29‰ in Porto de Mós). The highest value of 2.95‰ observed in the Figueira da Foz sector is recorded in the extremely base of bifrons Zone perhaps because the middle-upper part of levisoni Zone is not observed. Isotopic curves illustrated in Figures 4 to 7 allow the following stratigraphic observations:

MLLF, dated as polymorphum Zone, shows a clear positive trend in Coimbra and Figueira da Foz sectors. This trend is particularly well observed in the composite-section of Coimbra, where $\delta^{13}$C shows an important excursion in the middle part of this member, from values below 1‰ reaching at the top 2.6‰ (1% to 2.15‰ in Figueira da Foz).

The greatest Toarcian changes in $\delta^{13}$C occur between the TNL and the underlying and overlying units MLLF and MMLHH, showing that the TNL has the lowest values of the whole series. The boundary between MLLF and TNL shows an abrupt drop (from 2.6% to -0.87‰ in Coimbra; 2.15% to -0.25‰ in Figueira da Foz).

The MMLHH shows a large increase at the base (middle part of levisoni Zone) of this unit, followed by a continuous and gradual drop, towards the top of the unit (bifrons Zone), to carbon isotope values of 1.71‰ in Rabaçal, 1.79‰ in Coimbra and 1.11‰ in Porto de Mós.

Despite the differences in absolute carbon isotope values between each section observed in the MMLSB (uppermost bifrons Zone to bonarellii Zone), this interval clearly show a negative trend. The Coimbra sector shows an evolution from 1.92‰ to 0.48‰, a similar difference to the evolution occurred in Porto de Mós (2.54‰ to 1.47‰) and Rabaçal (2.52‰ to 1.7‰).

Although not analysed in detail, the MMLB shows at the base (speciosum Zone) a small increase in the $\delta^{13}$C, compared to the underlying unit. These values range between 1.19‰ in Coimbra, 1.77‰ in Rabaçal and 2.05‰ in Porto de Mós.

### Oxygen Isotopes

Oxygen isotopic results for the Lower and Middle Toarcian of LB show a wide range of values below -2.00‰, reaching values around -5.00‰. This large range and the negative values of $\delta^{18}$O are interpreted here to be best explained by a diagenetic overprint (early cementation or burial diagenesis), due to the known thermodependence of oxygen isotopes. For this reason, these data are not included in the aims and discussion of this work.

### Sequence stratigraphy

**Toarcian 2nd order Sequence: ST**

The marly limestone of the Toarcian succession in the LB comprises marl-limestone alternations, in some cases highly rhythmic and organised into hierarchical sequences. Two scales of sequences (2nd to 3rd order; Vail et al., 1991), equivalent to stratigraphic cycle terminology of Jacquin and De Graciansky (1998a,b), have been distinguished in the Toarcian series of the LB: 2nd order transgressive-regressive (T-R) facies cycles and 3rd order sequence cycles. The studied series are included in the 2nd order sequence ST (Duarte et al., 2004a,b). This sequence embraces the whole S. Gião Fm and the lowermost part of Póvoa da
Lomba Fm, varying between 75-80m in Tomar and around 280m thick in the Coimbra sector.

The base of ST, dated from the lowermost *polymorphum* Zone, corresponds to an abrupt flooding event, represented by marl accumulation in the whole basin. The marl dominance observed at the top of the *levisoni* Zone marks the peak transgression of the Toarcian 2nd order sequence (Duarte et al., 2004a). The Middle Toarcian-Lower Aalenian succession shows a regressive trend, ending ST with an upward increase of calcareous and bioclastic content. The upper discontinuity is dated as *opalinum* Zone and shows different sedimentary records in the basin (DA1 in Duarte et al., 2001).

**Toarcian 3rd order Sequences**

The Toarcian 2nd order sequence is subdivided into four 3rd order depositional sequences (ST1 to ST4; Figures 4 to 7), each one bounded by isochronous regional discontinuities (SBT1-SBT4; DT1-DT4 in Duarte, 1997), recognised over most of the LB (Figure 8). These sequence boundaries are dated of lowermost *polymorphum* Zone (SBT1), *polymorphum-levisoni* Zone boundary (SBT2), *intra-bifrons* Zone (SBT3) and uppermost *bonarelli* Zone (SBT4). Based on ammonite biostratigraphic data, and assuming ammonite zones of equal duration, the time span of each sequence agrees with the time range estimated by Vail et al. (1991) for the 3rd order depositional sequences (0.5 to 3 My).

**Sequence ST1 (polymorphum Zone):** In all the studied sections, the base of sequence ST is a clay-rich unit (MMLF), composed of decimetre- to metre-thick marl and centimetre-thick limestone alternations, very rich in benthic and nektonic macrofauna. Brachiopods (*Koninckella liasina, Nannirhynchia pyg...
moea, Liospiriferina), belemnites and ammonites are dominant and, in some levels, Zoophycos is very common, but benthic macrofauna decreases upward. This sequence increases in thickness towards north and west, reaching 20 m in Coimbra. The sedimentological features found across the polymorphum Zone of all the LB sectors suggest a clear deepening phase. ST1 is strongly asymmetric, and a large part of it is interpreted as a transgressive systems tract as a consequence of the increase of clay and the decrease of benthic macrofauna observed upwards. The maximum flooding surface probably corresponds to a thick limestone bed (0.50-0.80m), a condensed level of dactylioceratid ammonites (Dactylioceras semice-latun SIMPS), observed near the top of this sequence. This horizon is a good biologic event easily recognizable in the studied sections and in the major part of the LB.

Sequence ST2 (levisoni to bifrons Zones): One of the most marked sedimentary changes occurred in the LB during the Early and Middle Jurassic is observed between the polymorphum and levisoni Zones (Duarte 1997). This change shows different and particular sedimentological features in several points of the basin such as Peniche and Arrábida-Setúbal, related to tectonic control (Wright and Wilson 1984; Kullberg et al. 2001).

In the northern sector of the basin, the lowest part of ST2 shows 3-4 m of unfossiliferous brownish clays and marls. These facies, practically absent south of a line from Coimbra to Figueira da Foz, pass upwards to a thin succession (less than 10 m thick) of centimetre-thick alternations of marlstone, brownish claystone and micritic to microsparitic limestone (TNL). The large diversity of limestone facies includes silty to fine sandy limestones, with parallel and convolute lamination, and small-scale cross stratification (climbing ripples, wave ripples and low-angle cross bedding). The occurrence of the trace fossil Thalassinoides is a typical feature of the base of this unit. Integrating all the information about the base of this sequence it is possible to conclude that during the time equivalent to the base of levisoni Zone an important sea-level fall occurred, favoured by tectonic control. The shallowing is confirmed by the lithofacies nature of TNL, interpreted by Duarte and Soares (1993) and Duarte (1997) as a consequence of a tempestitic/turbiditic event. This tectonic phase is responsible for the western input of siliciclastics in the basin, particularly observed in Peniche (Wright & Wilson 1984; Duarte 1997).

The transgressive systems tract occurs in the whole studied area, through the decrease of siliciclastics and the increase of marls. This part of ST2, coincident with MMLHH, is composed of marl/limestone decimetre- to metre-thick alternations, particularly rich in rhynchonellid and terebratulid brachiopods (e. g. Soaresirhynchia bouchardi and Telothyris jauberti). This transgressive event ends before the progressive increase of calcareous components in the series. In sectors such as Coimbra and Figueira da Foz, this unit contains horizons rich in thin-shelled bivalve (Bositra sp.), pelagic evidence dated as the top of levisoni Zone. These horizons correspond to the maximum flooding interval of ST2 and to the peak transgression of the 2nd order sequence ST. These levels are overlain by thick aggrading/prograding package. In the whole basin, the series becomes progressively more calcareous evidencing a shallowing-upward evolution (highstand systems tract).

Sequence ST3 (bifrons Subzone – bonarelli/speciosum Zone transition): This sequence corresponds to
the MMLSB and it is characterized by marl/limestone alternations with siliceous sponge mud mounds (Duarte, 1997; Duarte et al., 2001). These bioherms are observed in all sectors with exception of Tomar, Peniche and Arrábida, occurring just to the top of ST3. The strong accumulation of marls (6 to 10 m thick) recorded at the base of this sequence in northern regions is interpreted as a transgressive systems tract, resulting from an increase of accommodation space. Laterally, these marls are equivalent in Porto de Mós to a siliceous sponge biostrrome condensed level that is interpreted by Duarte et al. (2001) as the ST3 maximum flooding interval. Overlying this surface is aggradational/progradational package belonging to the interval between gradata and the top of bonarellii/base of speciosum Zone. The regressive trend is well-recorded in some eastern sectors (Rabaçal-Alvaiázere) and particularly in Tomar and Peniche (Duarte, 1997). In the first case, the series becomes progressively more calcareous and bioclastic, ending the sequence with a rhynchonellid (Pseudogibbirhynchia sp.) bioclastic horizon. In Tomar, the bioclastic enrichment is evident in biosparites/grainstones with bivalves (Plagiostoma sp., Trigonia sp.), brachiopods (rhynchonellids and terebratulids), echinoid spines and ahermatipic corals (Figure 8). In Peniche, the marls that characterize the basin-scale uppermost bifrons Zone are overlain by
an oolitic-peloidal/marl succession. This prograding trend, observed in the Peniche section, is associated with a hemipelagic to outer submarine fan evolution (Wright and Wilson 1984).

**Base of Sequence ST4:** This sequence includes the units MMLB (last member of S. Gião Formation) and the base of Póvoa da Lomba Formation. It belongs to the *speciosum–opalimum* Zone interval. Similarly to the preceding ST3 sequence, ST4 starts with a very strong accumulation of marls and argillaceous limestone, particularly important in the northern sectors of the basin where MMLB can reach around 40 m thick. The base of MMLB is interpreted as a transgressive systems tract, as a consequence of an increase in accommodation space. The maximum flooding interval is represented by the massive occurrence of brachiopod-rich horizons (with *Soaresirhynchia renzi* and *Nannirhynchia cotteri*), well developed in several distal parts of the basin (Almeras, 1994).

**Discussion**

*$\delta^{13}C$ and Sequence Stratigraphy Interpretation*

It is known that $\delta^{13}C$ is a useful tool in stratigraphic analysis, particularly in the case of hemipelagic marl-limestone alternation successions. This is the case of the Lower and Middle Toarcian series of the LB, which also has an accurate ammonite biostratigraphy and displays 3rd order transgressive and regressive cycles throughout. The $\delta^{13}C$ composition of bulk carbonate reflects complex relations between ocean and atmosphere (e.g. Weissert, 1989; Weissert *et al.*, 1998; Kump and Arthur, 1999) at time of deposition, and probably diagenetic burial reactions, but variations of this signal in the Toarcian of LB also reflect clear changes in the depositional system. In fact, 3rd order sequence boundaries are always accompanied by important, sometimes large variations, in $\delta^{13}C$ composition. These observations indicate a control-
ling mechanism of primary nature and clearly associated with transgressive or regressive phases. Transgressive Phases

According to the interpretation above, ST1 is truncated at the top, as a consequence of an input of siliciclastics in the series, generated by tectonics (Duarte et al., 2004a). This tectonic phase, dated as lowermost levisoni Zone, interrupts a large 3rd-order transgressive event that characterizes ST1. This transgressive trend is well correlated by the positive excursion observed in δ13C at the top of this unit (top of MLLF), reaching very similar values in the distal sectors of Coimbra and Figueira da Foz (Figures 6 and 7).

Both ST2 and ST3 show a relatively short transgressive phase, more evident in the case of ST3. Considering the first sequence (ST2), the base of MMLHH (middle-upper part of levisoni Zone), that is interpreted as a 3rd-order deepening phase, shows a strong positive isotopic shift, reaching values of 4.29‰ in Porto de Mós and 4.02‰ in Coimbra; this excursion shows a large change of 4‰ in a time duration of only about 0.5 My.

The large increase of δ13C observed in the extreme (transgressive) base of ST3 at Porto de Mós (Figure 5) is good evidence of this relation between carbon-isotope variation and sequence evolution.

Regressive Phases

In contrast to ST1, ST2 and ST3 are both characterized by strong development of regressive phases. In both cases δ13C curves show a clear and significant decrease along the middle-upper part of each 3rd-order sequence but with different magnitudes. Comparing the results observed in Coimbra, Rabaçal and Porto de Mós across ST2, the amplitude of decrease is very similar in the first two areas (2.23‰, 2.36‰ respectively) while in Porto de Mós it is greater than 3.16‰, reaching 1.11‰ (Figures 4 to 6). In fact, the depositional contrast between the top of ST2 and base of ST3 is better expressed in the Porto de Mós section (discussion in Duarte et al. 2001). The regressive trend of ST3 correlates to a slight negative trend in the δ13C record: from 2.52‰ to 1.7‰ in Rabaçal, 2.54‰ to 1.47‰ in Porto de Mós and 1.89‰ to 0.48‰ in Coimbra (Figures 4 to 6). In this last sector, a negative excursion is observed at the top of ST3.

On the other hand, the lowstand deposits that characterizes the base of ST2 (TNL) are correlatable with an abrupt decrease of carbon-isotope evolution, reaching values around 0‰ (-0.87‰ in Coimbra; -0.25‰ in Figueira da Foz).

Lower Toarcian Global Events

The Early Toarcian is known as an important time of global scale events. It is argued to be an important mass extinction episode (Little and Benton, 1995; Pálfy and Smith, 2000), associated with oceanic anoxia (Jenkyns, 1988). More recently, these events have been related with the austral Karoo-Ferrar flood basalt episode (Pálfy and Smith, 2000; Jones and Jenkyns, 2001). The Lower Toarcian (around polymorphum/levisoni boundary) deposits of the LB show the most important sedimentological changes that occurred in the whole Early Jurassic (Duarte, 1997; Duarte et al., 2004a), connected with a strong perturbation in the biological record (Mouterde and Ruget 1984; Gahr, 2002, 2005; Perilli and Duarte, 2006), but without the development of anoxic facies. In fact, in Portugal, the extreme base of levisoni Zone is marked by a strong reduction of benthic and nektonic macrofauna record. As demonstrated above, the negative and positive carbon-isotope excursions are clearly associated with 3rd order regressive and transgressive phases, respectively. However, carbon-isotope curves observed in the studied sections seem to reflect also the main global isotopic events (Figure 9).

a) The most important δ13C positive excursion (magnitude up to 4‰) observed in the middle-upper part of levisoni Zone of LB is more or less associated with 2nd order sequence maximum flooding phase. Despite the absence of organic carbon-rich sediments in the studied sections (in Portugal, only the levisoni Zone of the Peniche section shows high TOC values; Duarte et al. 2005), this isotopic positive excursion is well observed in several European and Tethyan basins (Jiménez et al., 1996; Jenkyns et al., 2002). This event corresponds to a widespread deepening phase, illustrated in several eustatic curves (Hallam, 1988; 2001; Haq et al., 1988; Hardenbol et al., 1998), and seems to be associated to the maximum flooding interval of 2nd order transgressive-regressive facies cycles observed in some Western European basins (e.g. Paris Basin, in De Graciansky et al., 1998b).

The gradual decrease of δ13C up to the bifrons Zone is consistent with the trends observed in neighbouring basins of Iberia and Central Europe (Jenkyns and Clayton, 1986, 1997; Jenkyns et al., 1991, 2002; Jiménez et al., 1996; Rosales et al., 2001).

b) The strong negative carbon isotopic excursion recorded at the base of levisoni Zone (TNL) is typical of the event that characterizes marine carbonates and
organic matter dated as uppermost *tenuicostatum (polymorphum)* Zone and the lower part of *falciferum (levisoni)* Zone of some Central European basins (Jiménez et al., 1996; Hesselbo et al., 2000; Schouten et al., 2000; Beerling et al., 2002; Jenkyns et al., 2002; Van de Schootbrugge et al., 2005). Some authors such as Hesselbo et al. (2000) and Beerling et al. (2002) suggested that this rapid event was the result of large methane gas-hydrate dissociation. Others, such as Schouten et al. (2000) and Röhl et al. (2001) considered recycling and re-utilization of $^{12}$C from stratified anoxic bottom waters as a possible mechanism. However, if we compare the known $\delta^{13}$C curves in different domains of Europe we conclude that this negative excursion is not so generalised as the following positive excursion. In fact, it is not recognised in some neighbouring sectors such as the Basque-Cantabrian Basin (Rosales et al., 2001).

Independently of the possible global mechanisms, the lowstand deposits observed in the LB may be justified by the input of $^{12}$C in the oceanic carbon reservoir that could have been released by erosion and weathering of continental material. Contrary to the observation of widespread black shale deposition, the Lower Toarcian (both *polymorphum* and *levisoni* Zones) of the four LB studied sections shows low values of organic matter (TOC values below 0.5%; Duarte et al., 2005).

Conclusions

The Toarcian successions in the central and northern part of the LB are dominated by hemipelagic marl/limestone alternations, rich in benthic and nektonic macrofauna. The Lower and Middle Toarcian series are organised into three 3rd order depositional sequences (ST1, ST2 and ST3), bounded by four isochronous regional discontinuities, well dated by ammonites: at base of *polymorphum* Zone (SBT1), across the *polymorphum/levisoni* Zone boundary

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**Fig. 9 – Correlation between 3rd order depositional sequences and $\delta^{13}$C evolution in the Toarcian of Lusitanian Basin.** Carbon-isotope events in carbonate of several neighbouring basins: 1 - Subbetic Cordillera (Jiménez et al., 1996); 2 - Basque-Cantabrian Basin (Rosales et al., 2001); 3 - Yorkshire (McArthur et al., 2000); 4 - Mochras well in Wales (Jenkyns et al., 2002; Van de Schootbrugge et al., 2005).

**Fig. 9 - Correlación entre las series deposicionales de 3° orden y la evolución del $\delta^{13}$C en el Toarcense de la Cuenca Lusitánica.** Eventos de los isotopos de carbono en los carbonatos de varias cuencas vecinas: 1.- Cordillera Subbética (Jiménez et al., 1996); 2.- Cuenca Vasconcántabra (Rosales et al., 2001); 3.- Yorkshire (McArthur et al., 2000); 4.- Pozo Mochras, en Gates (Jenkyns et al., 2002; Van de Schootbrugge et al., 2005).
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(SBT2), within the bifrons Subzone (SBT3) and uppermost bonarelli Zone (SBT4), respectively. The most significant sedimentological change is observed at the ST1/ST2 boundary, where the facies of the base of levisoni Zone are interpreted as the result of a strong shallowing (lowstand).

Isotopic curves obtained in four reference sections of the LB, reveal that δ¹³C is an important tool in sequence stratigraphic interpretation. The first main conclusion of the observed variations is that the link with the sequence evolution should indicate a mechanism of a primary nature. The δ¹³C curves clearly show remarkable changes at the boundaries between the several 3rd order depositional sequences. The large positive fluctuations correspond to transgressive periods. With the exception of ST1, the uppermost parts of ST2 and ST3 show gradual negative shifts, associated with regressive phases. Two main positive δ¹³C excursions are observed at the top of polymorphum Zone and in the middle-upper part of levisoni Zone, both peaks associated with maximum flooding intervals. Between these two peaks, the minimum value (negative excursion) is observed at the base of levisoni Zone, associated with shallow detrital facies deposits.

Besides, the δ¹³C evolution observed in the Lower and Middle Toarcian of LB shows the global carbon shift recognised in other Tethyan and North-European sectors.

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